



An Integrated Rate Methodology (IRM) for Estimating Terrestrial Water and Carbon Fluxes

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Carbon-Water Coupling

Ta



Hot desert



Tropical forest



Deciduous forest

Boreal forest



Tundra

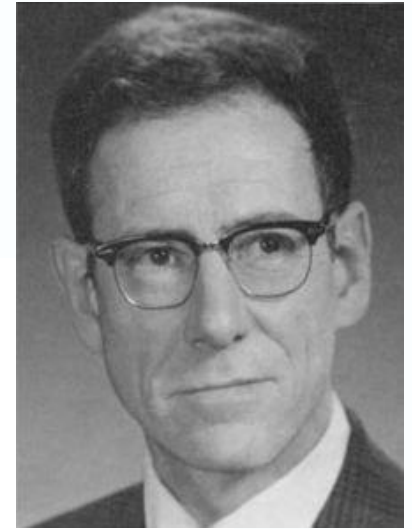


- Key controls of terrestrial ecosystem systems?
- Interactions/feedbacks over different scales?

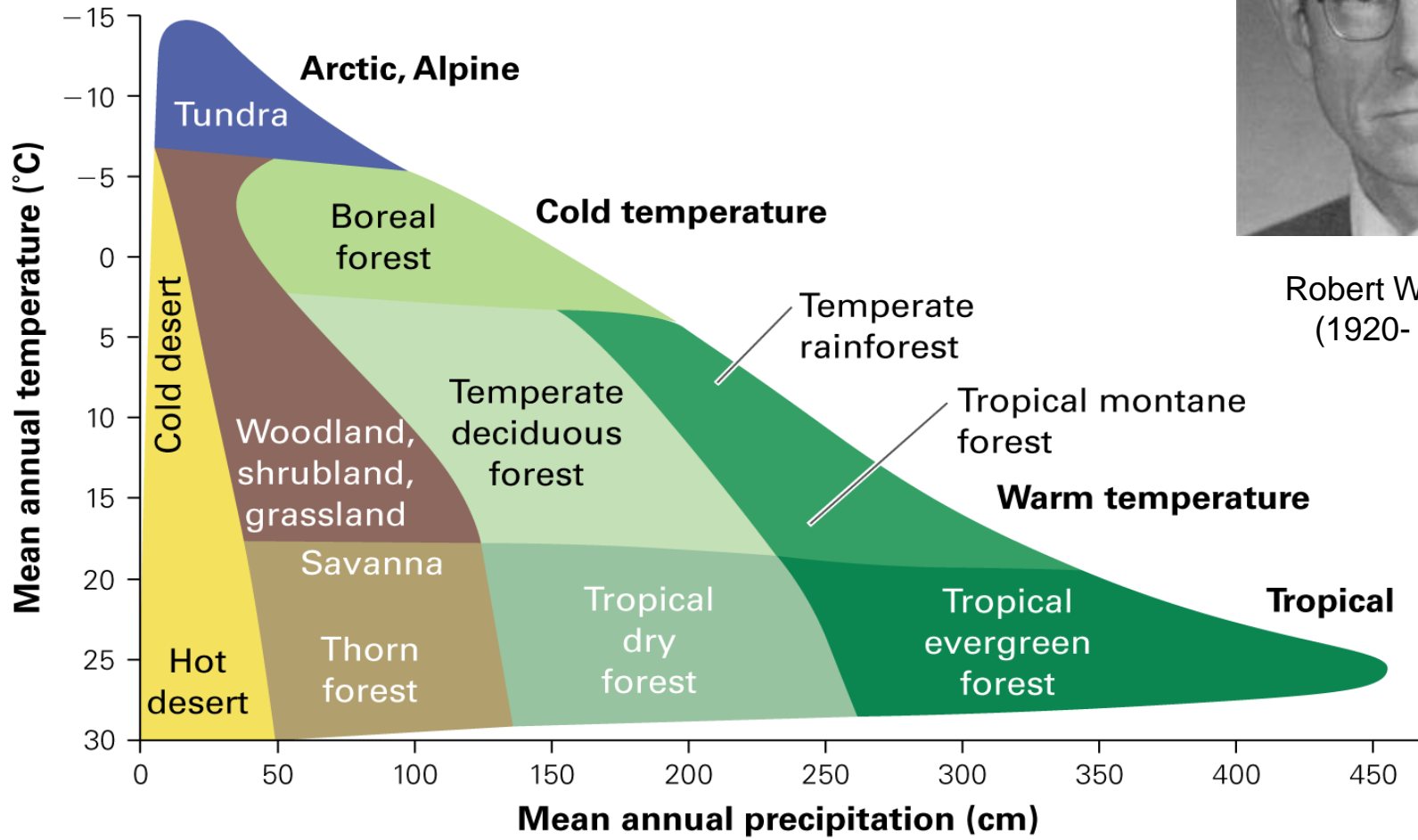
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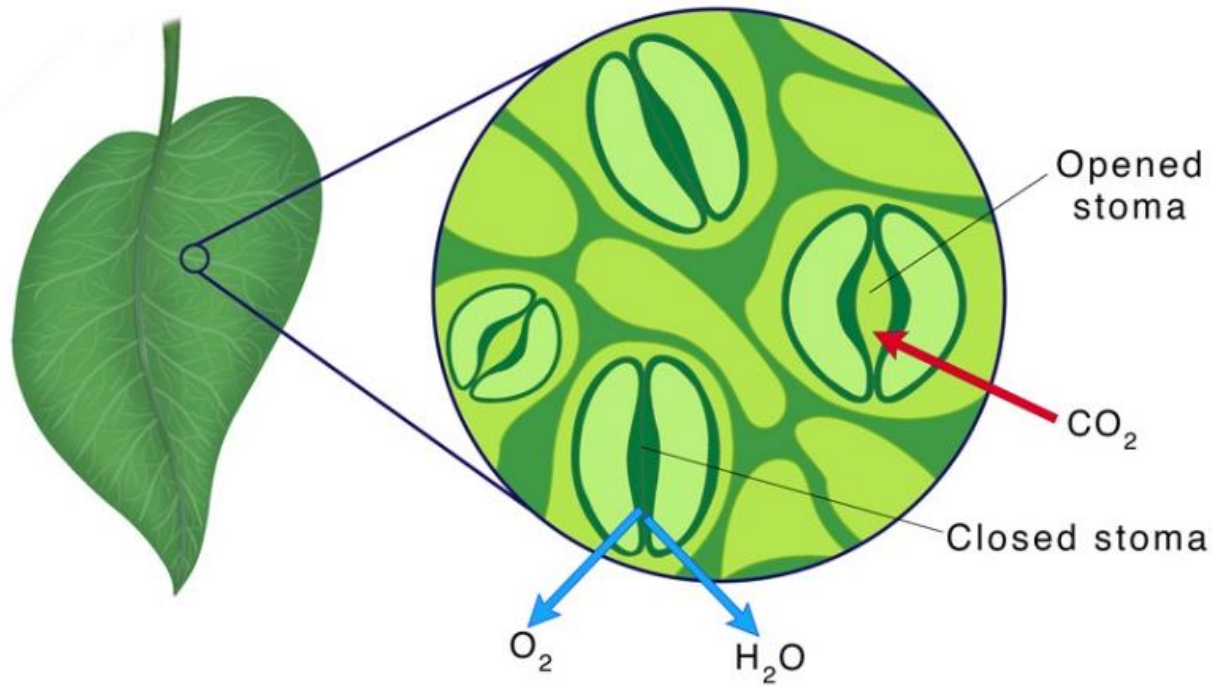
Biome and climate relationship: Whittaker's diagram



Robert Whittaker
(1920- 1980)



Carbon and water cycles are intrinsically coupled



Since carbon uptake and water loss occur through stomata, photosynthesis and transpiration both decline with stomatal closure.

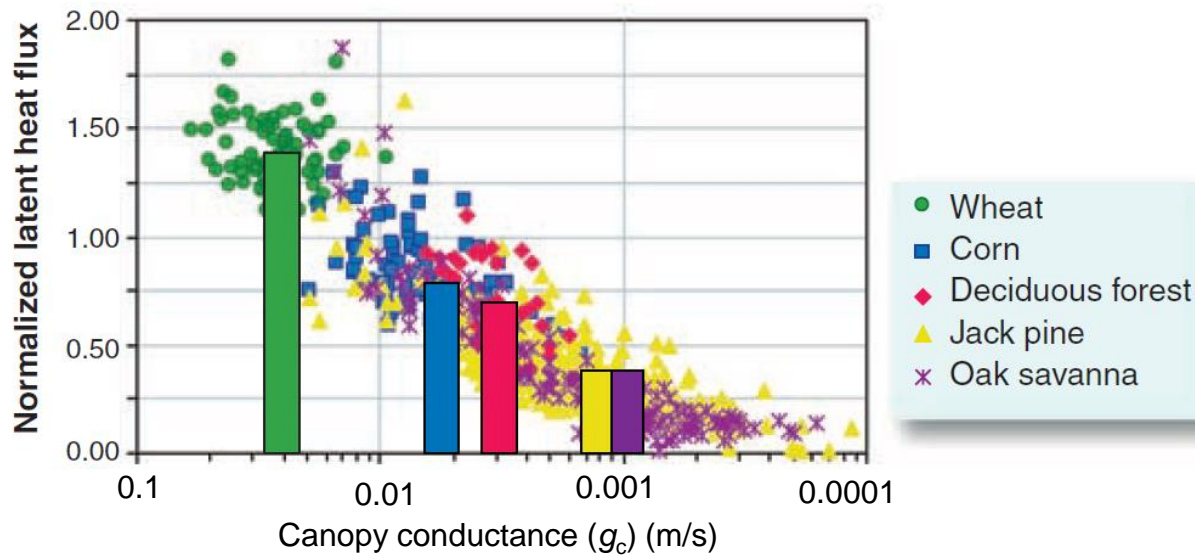
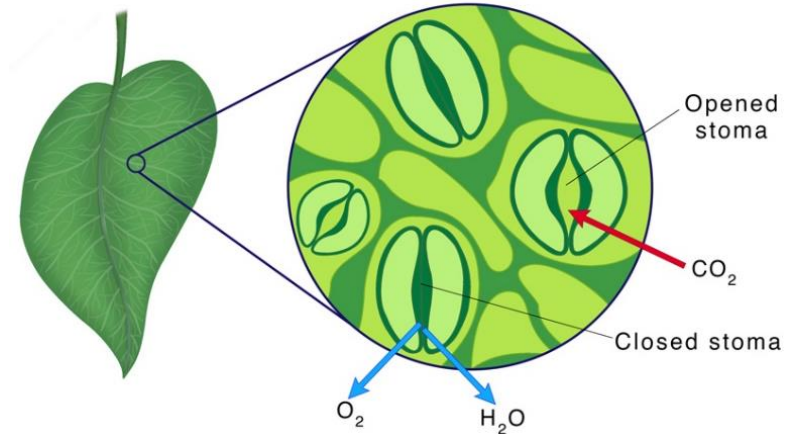


Transpiration and stomatal control

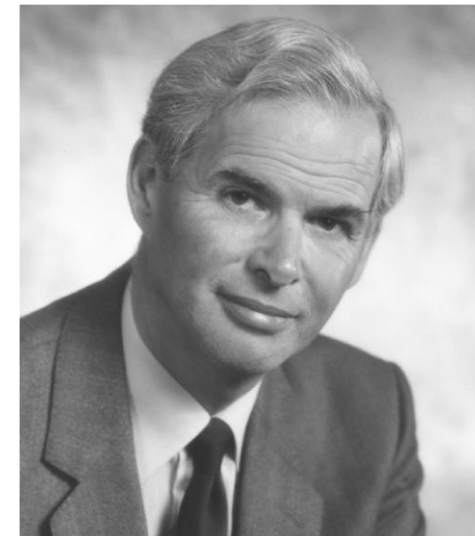
$$E = \rho g_c (q_s^* - q_s)$$

Penman-Monteith equation

$$E = \frac{\varepsilon A + (\rho c_p / \gamma) g_a D_a}{\varepsilon + 1 + g_a / g_c}$$



Bonan (2008)



John Monteith
(1929-2012)



Effect of stomatal control: from leaf to region

$$E = \Omega_{ci} E_{eq} + (1 - \Omega_{ci}) E_{imp} \quad \Omega_{ci} = (\varepsilon + 1) / (\varepsilon + 1 + g_a / g_c)$$

Ω_{ci} is the decoupling factor between vegetation and atmosphere.

$$\frac{dE}{E} = (1 - \Omega_{ci}) \frac{dg_c}{g_c}$$

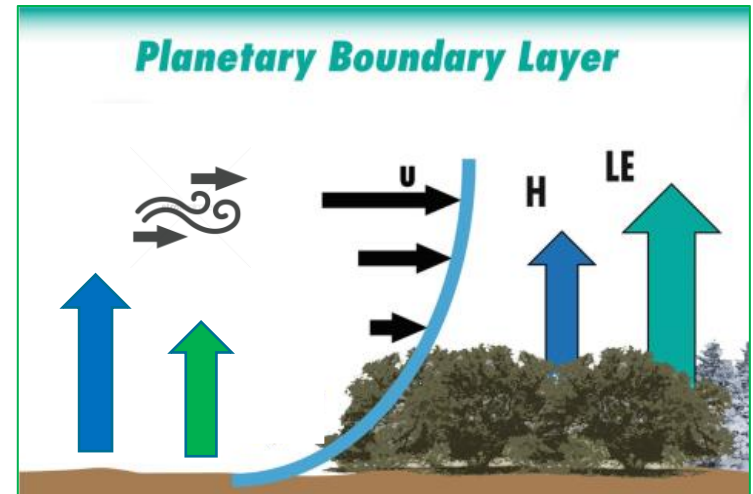
Ω_{ci} increases with spatial scale, E is less controlled by stomatal conductance.



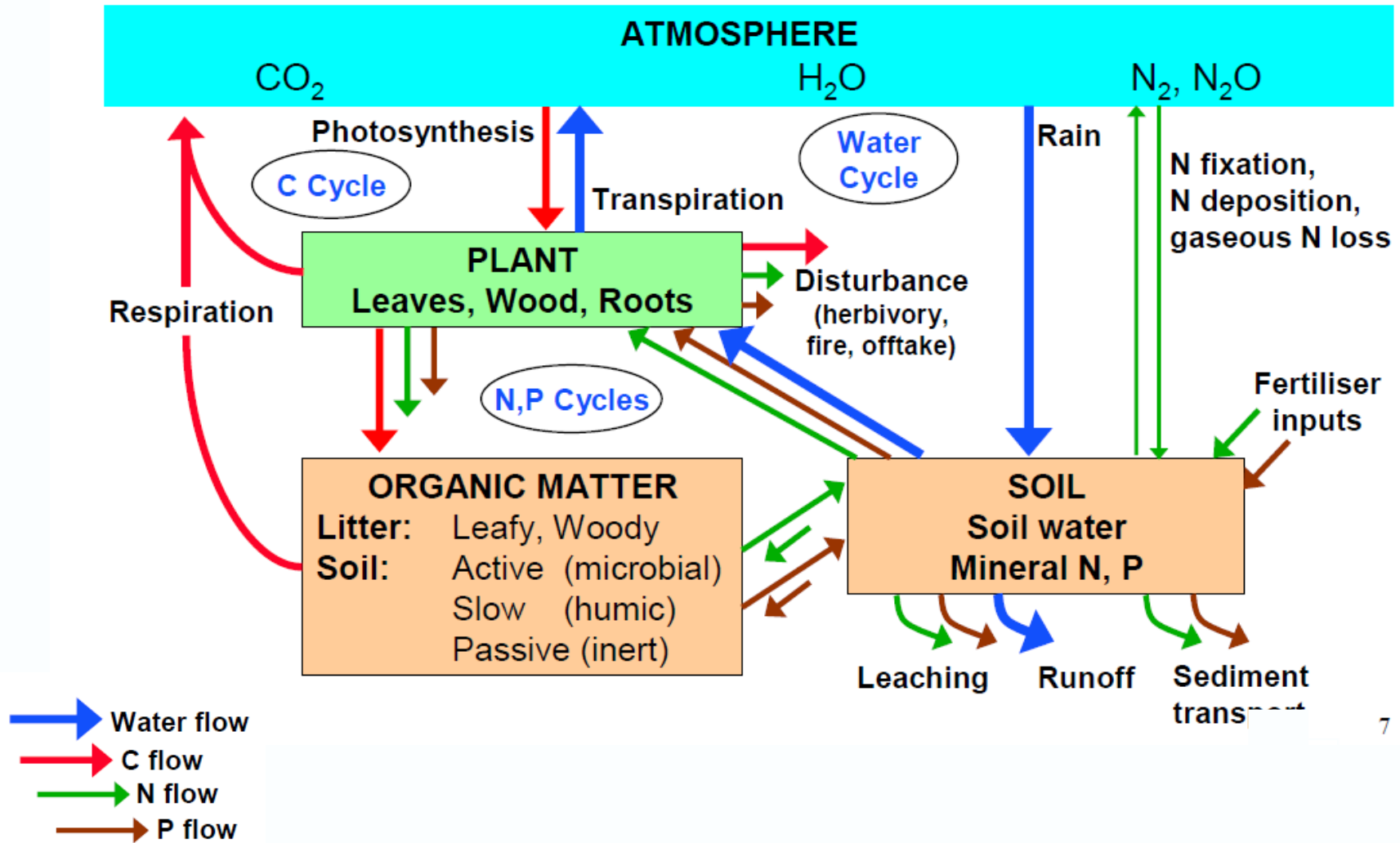
equal-proportional change

depends on Ω

much smaller change

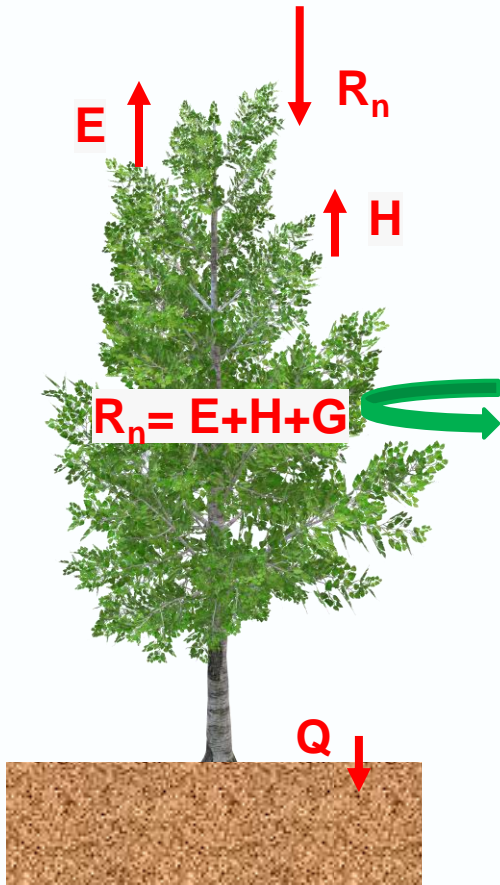


Linked terrestrial water and carbon cycles

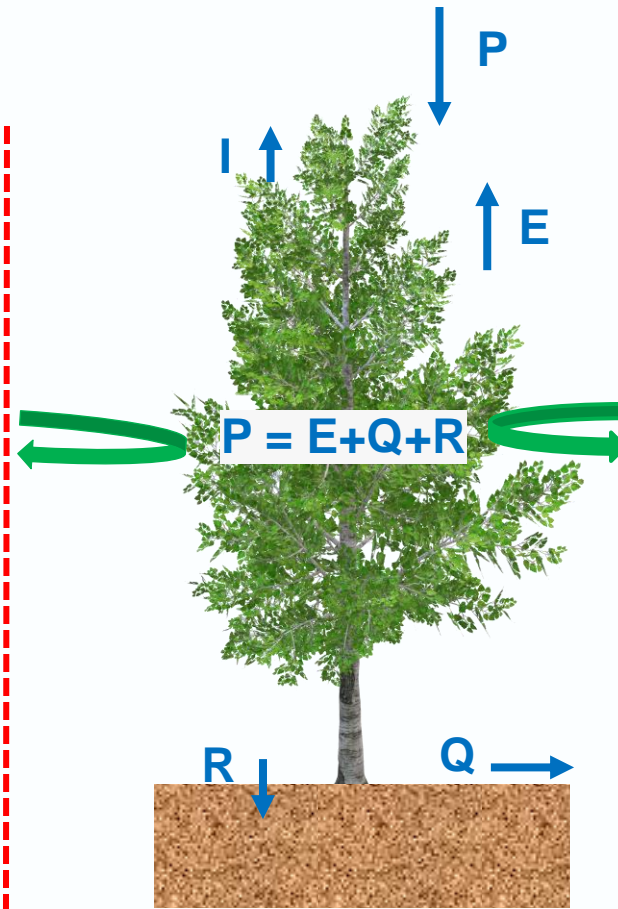


Energy-Water-Carbon Coupling

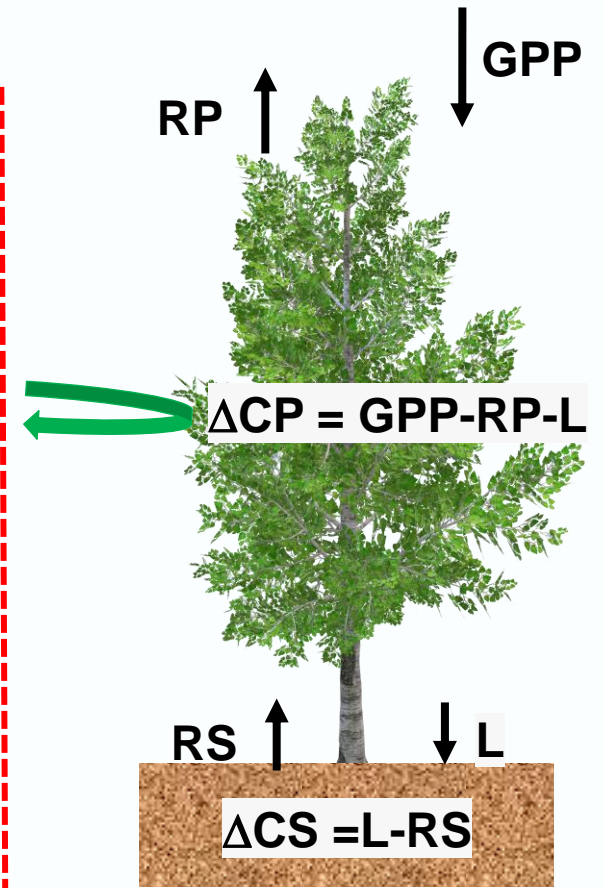
Energy Balance



Water Balance



Carbon Balance



Dynamics of water-energy-carbon fluxes

Water balance: $\frac{dS(t)}{dt} = P(t) - E(t) - Q(t) - R(t)$

Carbon balance: $\frac{dC_p(t)}{dt} = GPP(t) - RP(t) - L(t)$

Energy balance: $\frac{dW(t)}{dt} = R_n(t) - E(t) - H(t) - F_p(t)$

We must consider the balance in the process representation so that model consistency, reliability and accuracy can be achieved.



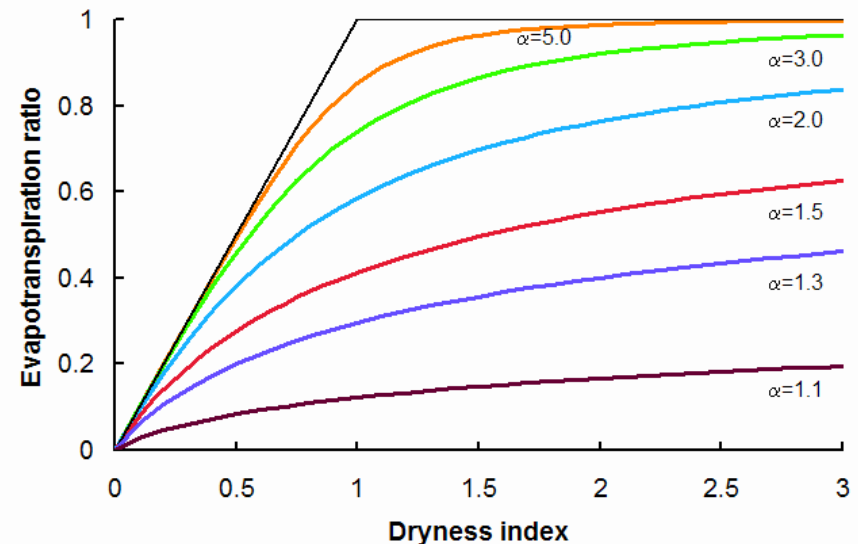
Theoretical framework: long-term average evaporation

Demand = Potential evaporation (E_0)

Supply = Precipitation (P)

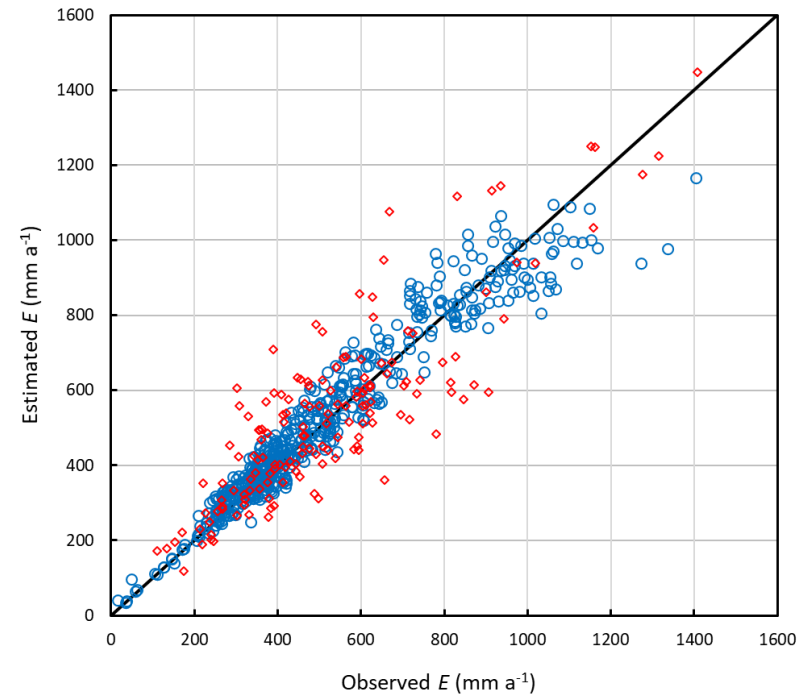
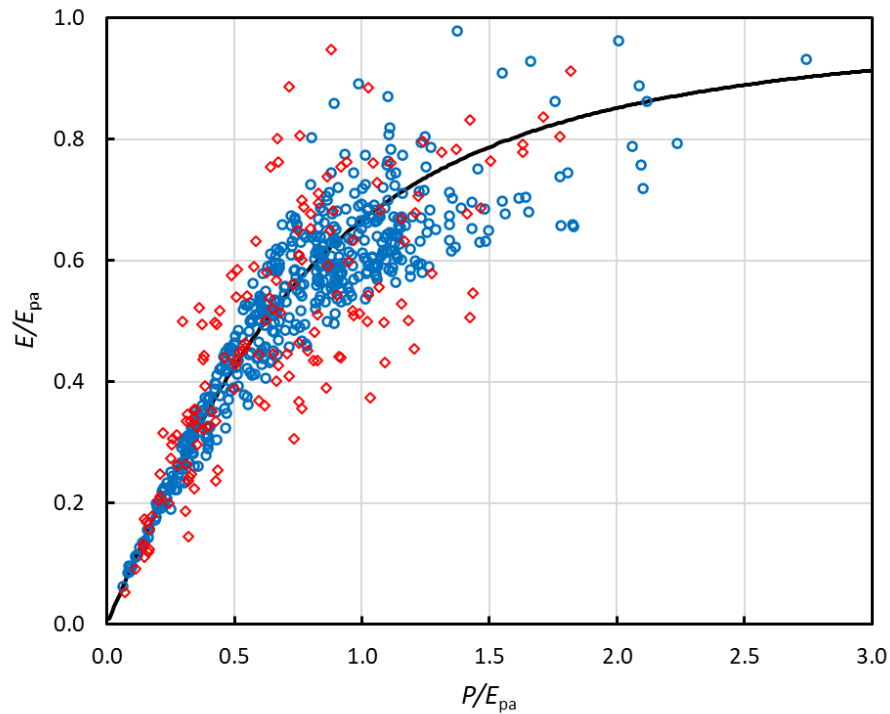
$$\frac{\partial E}{\partial P} = f(E_0 - E, P)$$

$$\frac{E}{P} = 1 + \frac{E_0}{P} - \left[1 + \left(\frac{E_0}{P} \right)^w \right]^{1/w}$$

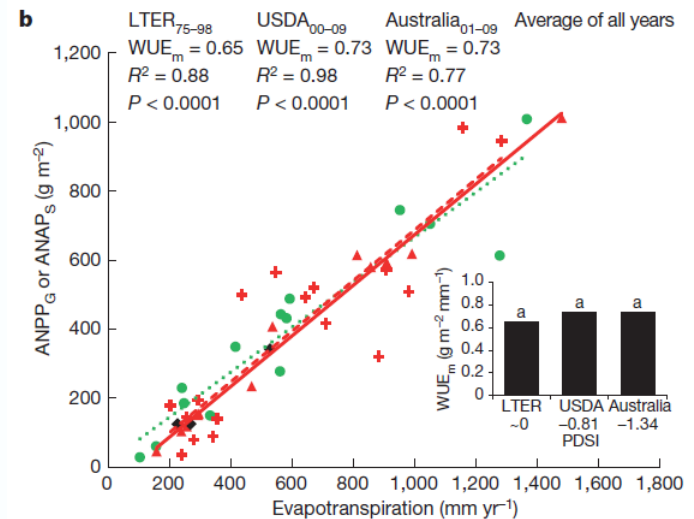
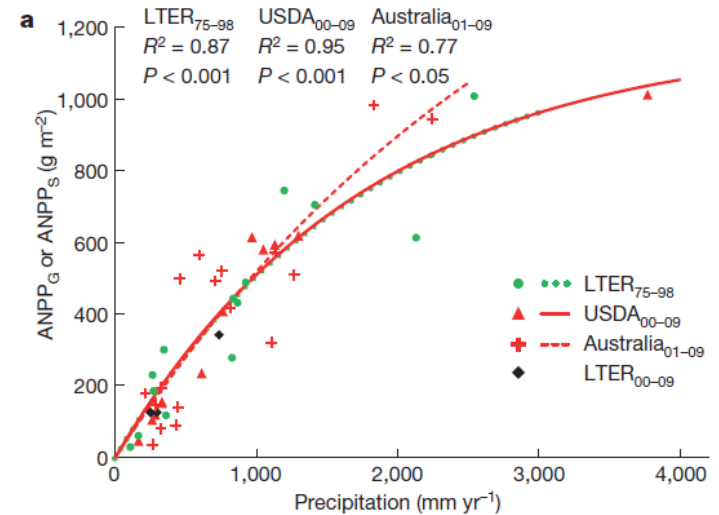
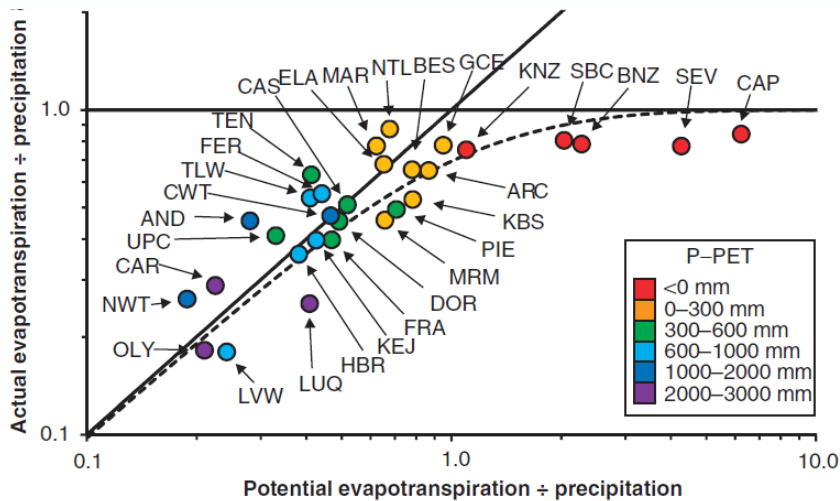
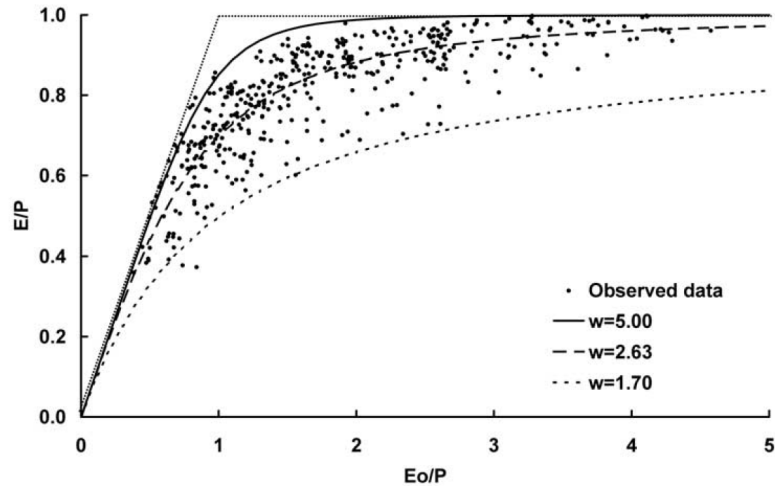


Theoretical framework: long-term average evaporation

$$\frac{E}{E_{pa}} = 1 + \frac{P}{E_{pa}} - \left[1 + \left(\frac{P}{E_{pa}} \right)^w \right]^{1/w}$$



Budyko-like equation for carbon?

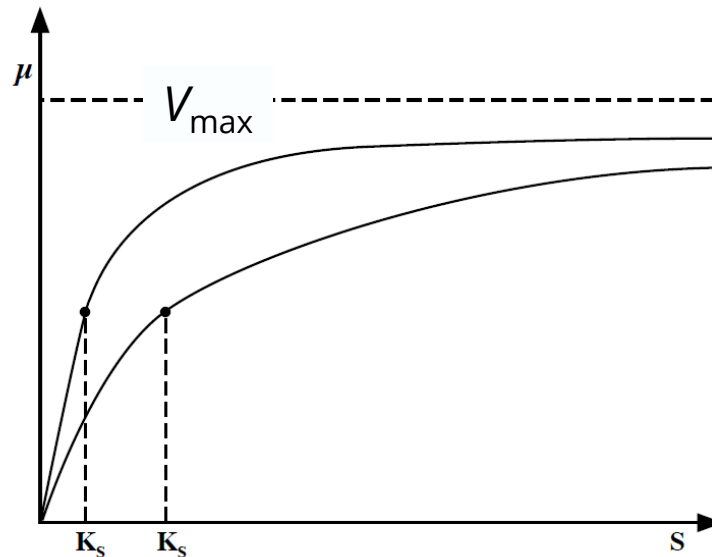


The Michaelis–Menten Equation

In biochemistry, the Michaelis–Menten Equation describes enzyme kinetics:

$$\mu = V_{max} \frac{S}{K_s + S}$$

Where μ is the reaction rate, and V_{max} is maximum reaction rate, S is the concentration of a substrate, K_s is the concentration of the substrate at which the reaction rate is half of V_{max} . K_s controls how fast V_{max} is approached.



Leonor Michaelis
(1875 -1949)



Maud Menten
(1879 -1960)



Generalized Michaelis–Menten Equation

To generalize a single substrate system to an n-substrate systems, a ratio form of the Michaelis–Menten equation is considered:

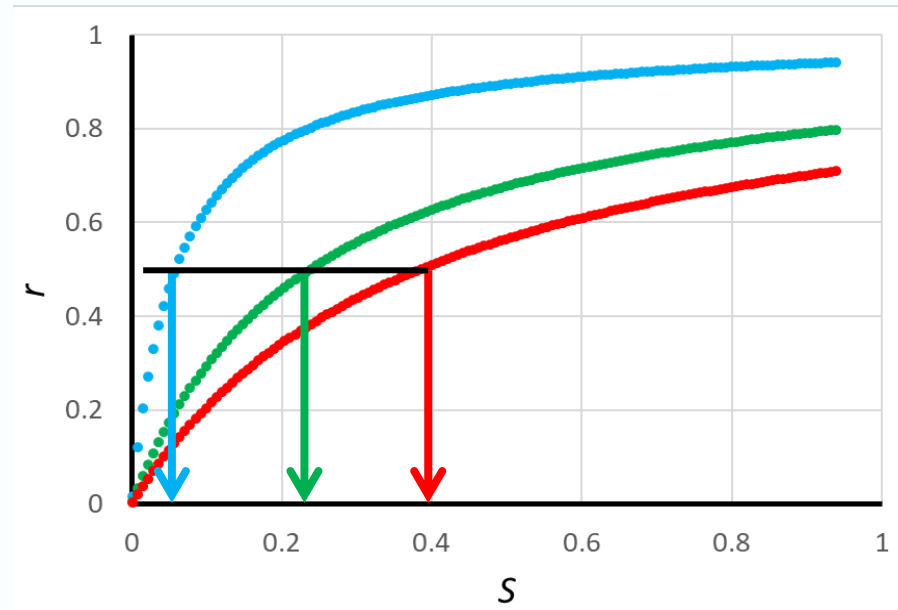
$$r = \frac{\mu}{V_{max}} = \frac{1}{1 + K_s/S}$$

The characteristics of this single substrate system:

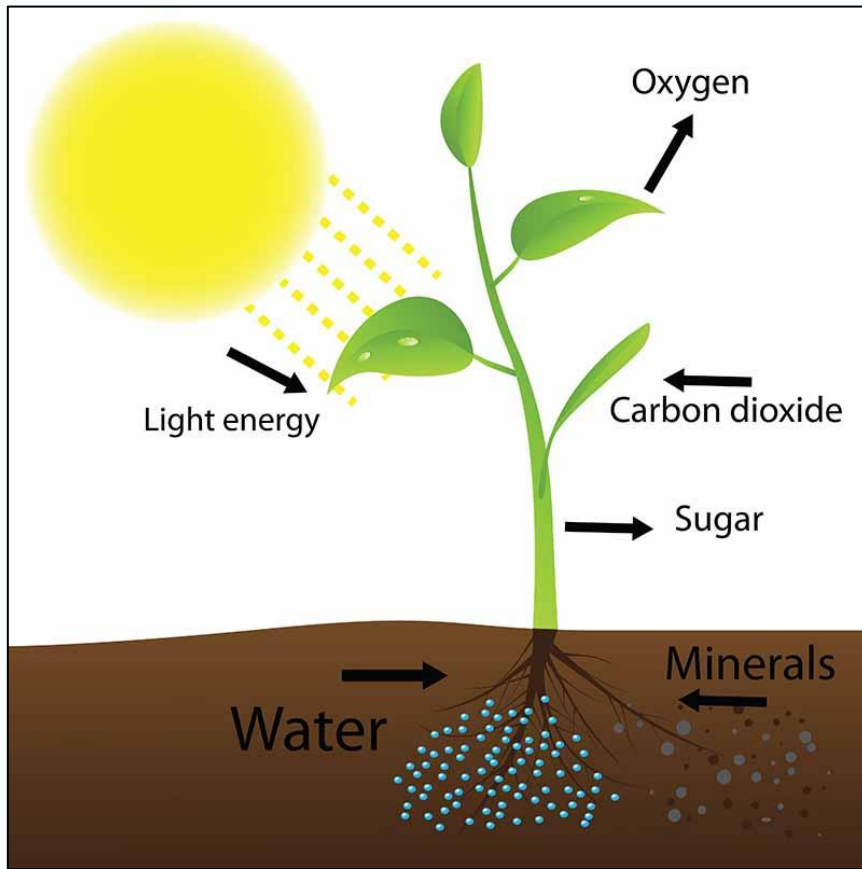
- $r = 0$ when $S = 0$
- $r = 1$ when $S \rightarrow \infty$
- $r = 1/2$ when $S = K_s$

Generalized Michaelis–Menten equation:

$$r_n = \frac{1}{1 + \sum_{i=1}^n (K_{si}/S_i)}$$



Plant growth is fundamentally a function of available light, water, and nutrients (Wu et al., 1994).



$$A = A_{max} \left[\frac{1 + W_H + W_N}{\frac{1}{m_L x_L} + \frac{W_H}{x_H} + \frac{W_N}{x_N}} \right]$$

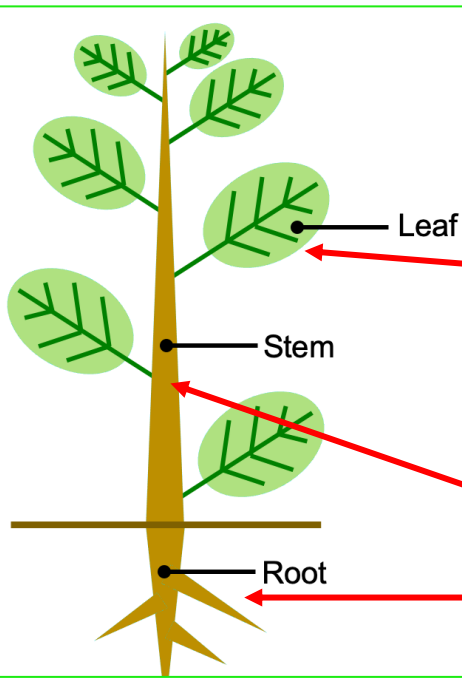
- light: x_L
- water: x_H
- nutrients: x_N

Modelling plant growth using IRM

$$A = A_{max} \left[\frac{1 + W_H + W_N}{\frac{1}{m_L x_L} + \frac{W_H}{x_H} + \frac{W_N}{x_N}} \right]$$

where W_H and W_N are the weightings of water relative to light and nutrients, x_L , x_H , and x_N are the relative resources availabilities for light, water, and nutrient respectively, and m_L is the modifier of light availability due to temperature.

The simulated carbon (A) is partitioned to leaves, stems, and roots:



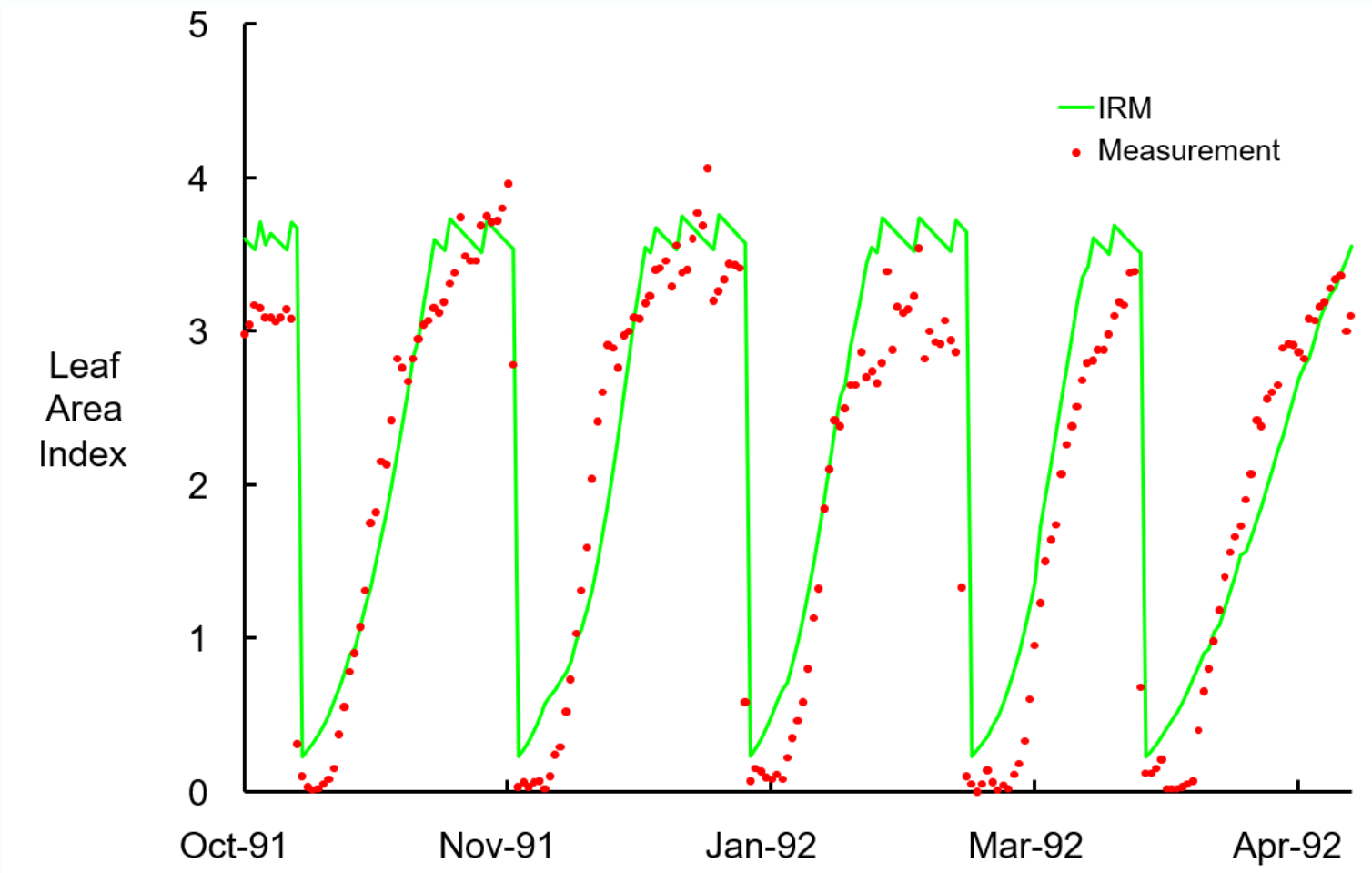
Leaf carbon

$$\Delta C_L = n_L Y_L (A - C_L R_L) - C_L M_L$$

Stem and root carbon

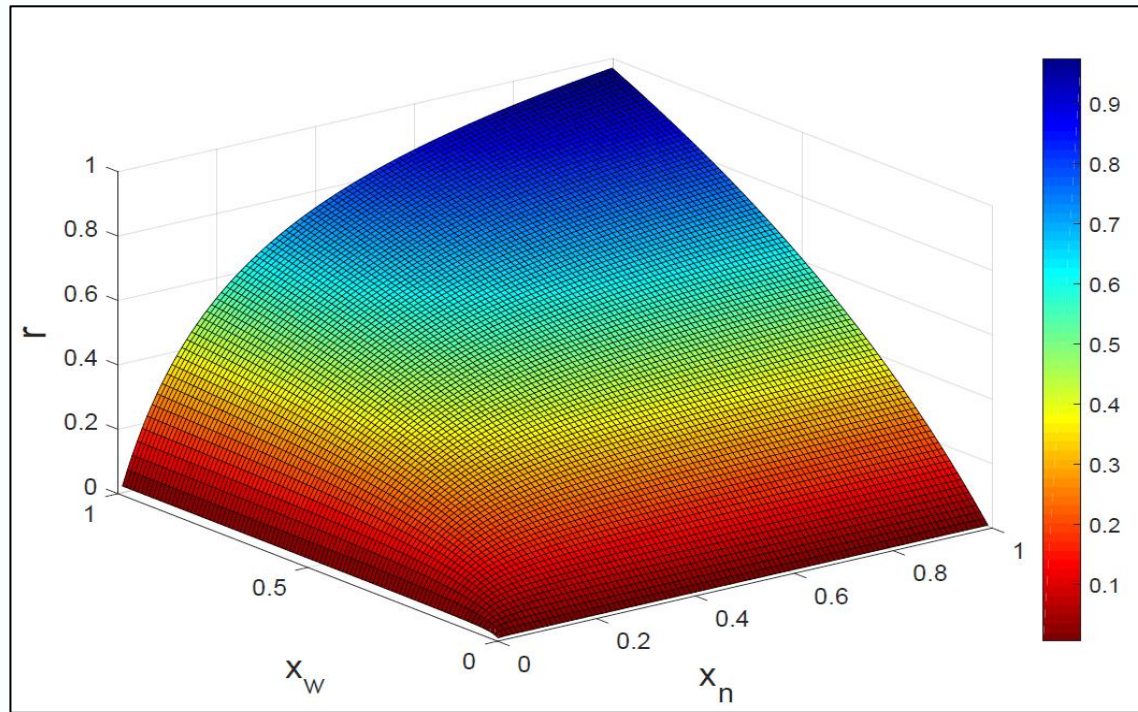
$$\Delta C_{S,R} = n_{S,R} Y_{S,R} \{ (A - C_L R_L) - C_{S,L} R_{S,R} \} - C_{S,R} M_{S,R}$$

Modelling plant growth using IRM



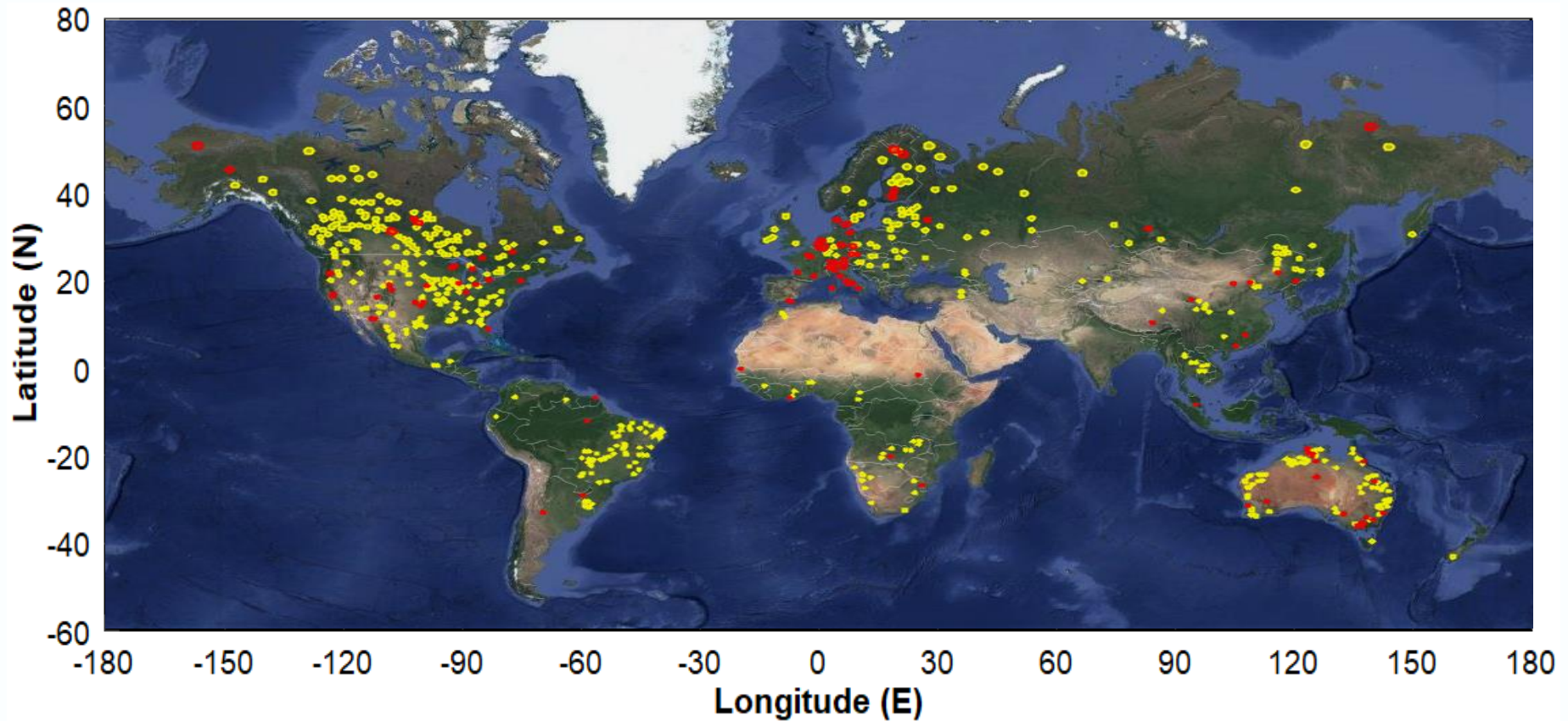
Integrated Rate Methodology (IRM) for mean annual ET and ANPP

$$Y = Y_{max} \left[\frac{1 + W_H + W_N}{\frac{1}{m_L x_L} + \frac{W_H}{x_H} + \frac{W_N}{x_N}} \right]$$



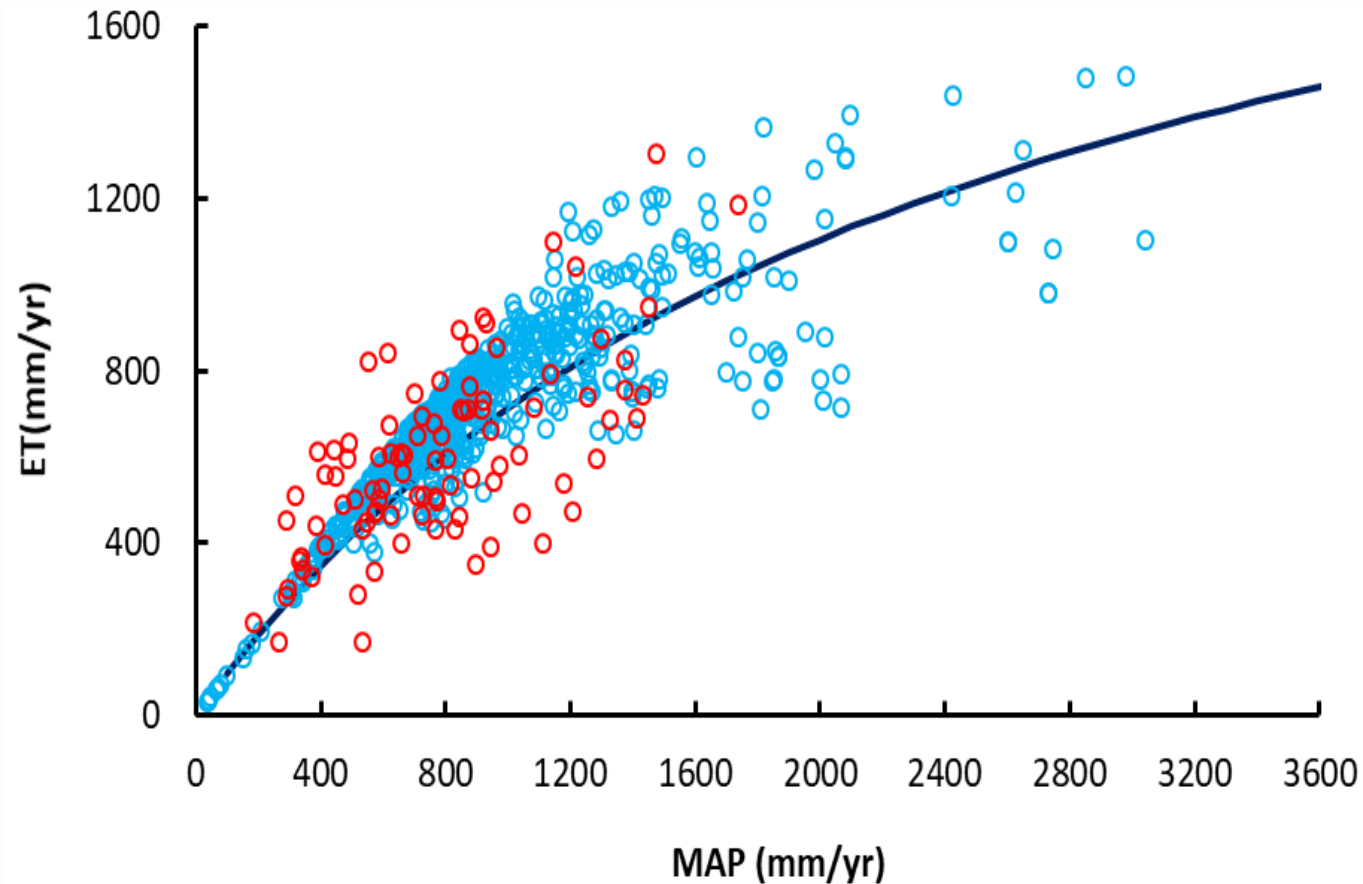
Global ET data:

- *Global catchment water balance* ($n=524$)
- *Global flux sites* ($n= 156$)

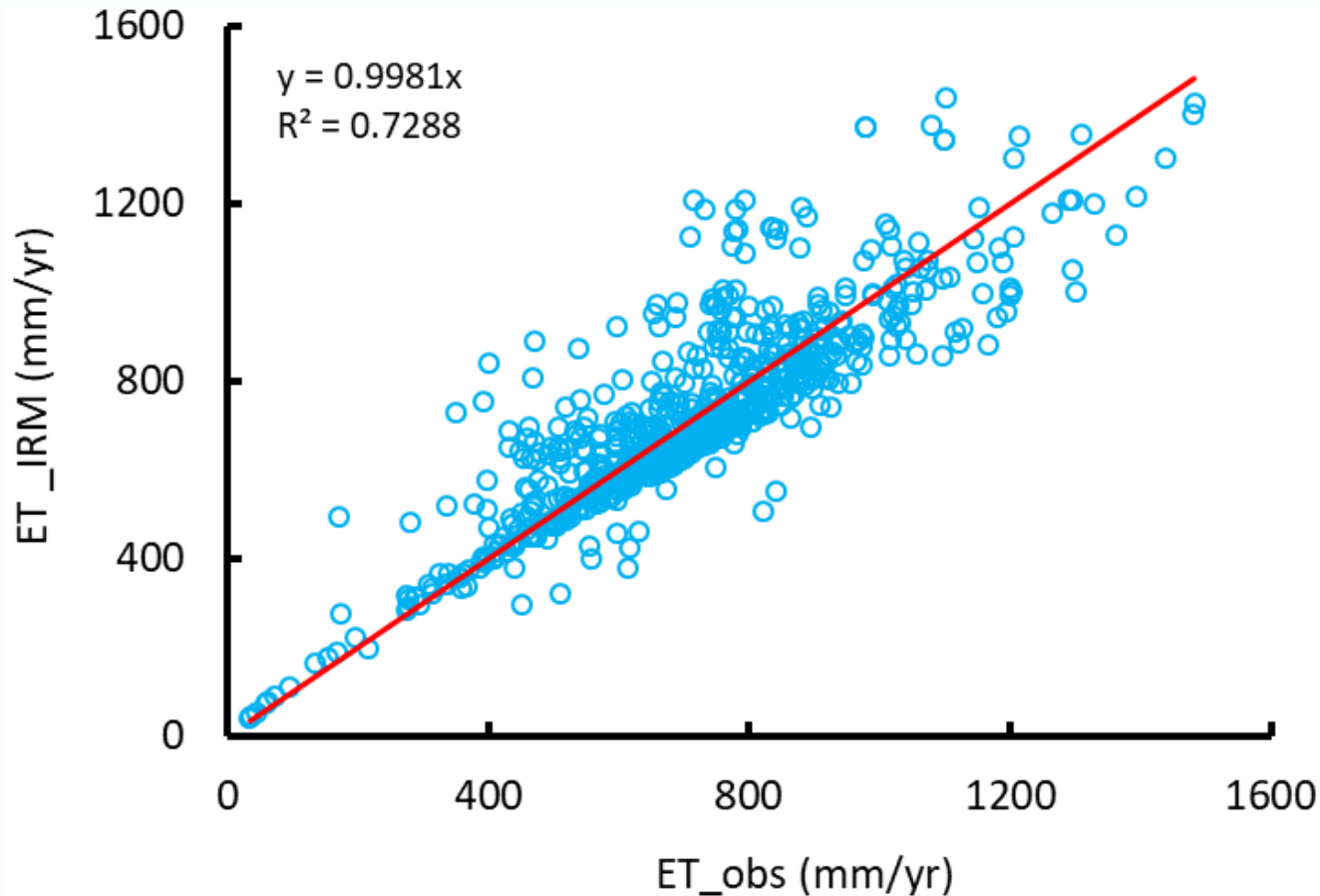


Estimation of mean annual evapotranspiration using IRM

Strong relationship between mean annual precipitation (MAP) and ET



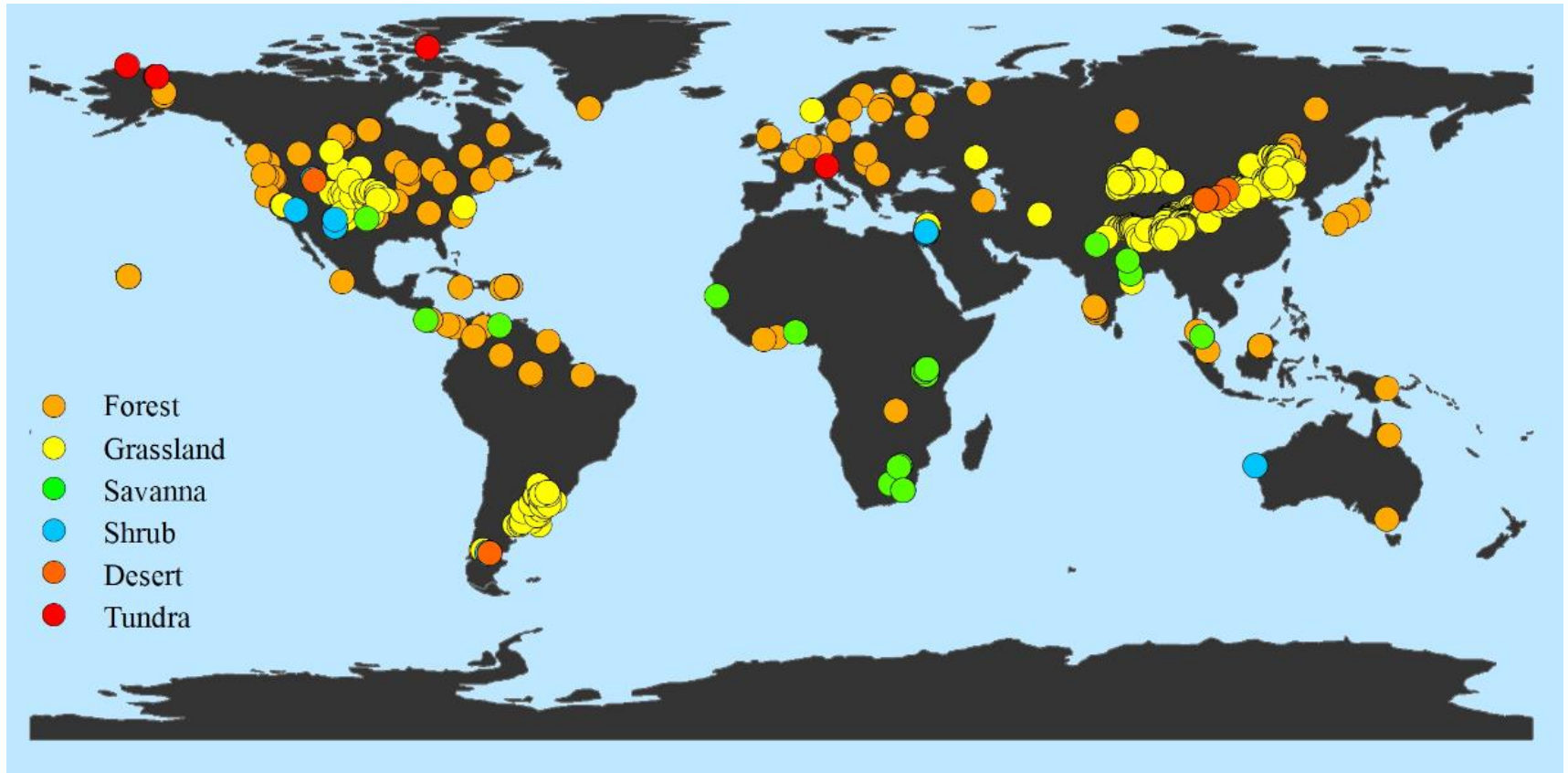
Estimation of mean annual evapotranspiration using IRM



➤ IRM can provide accurate estimates of mean annual ET

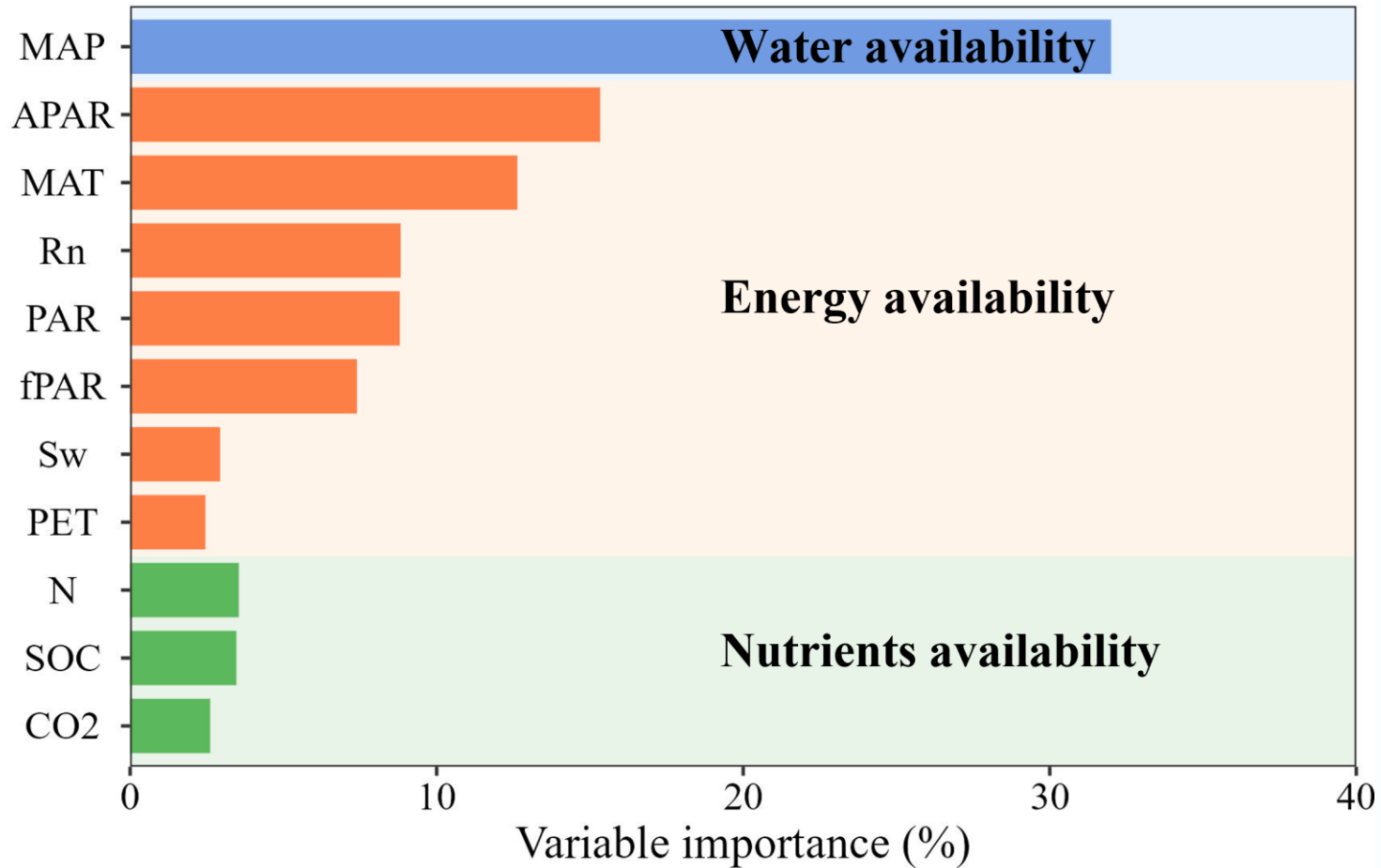


Global ANPP data



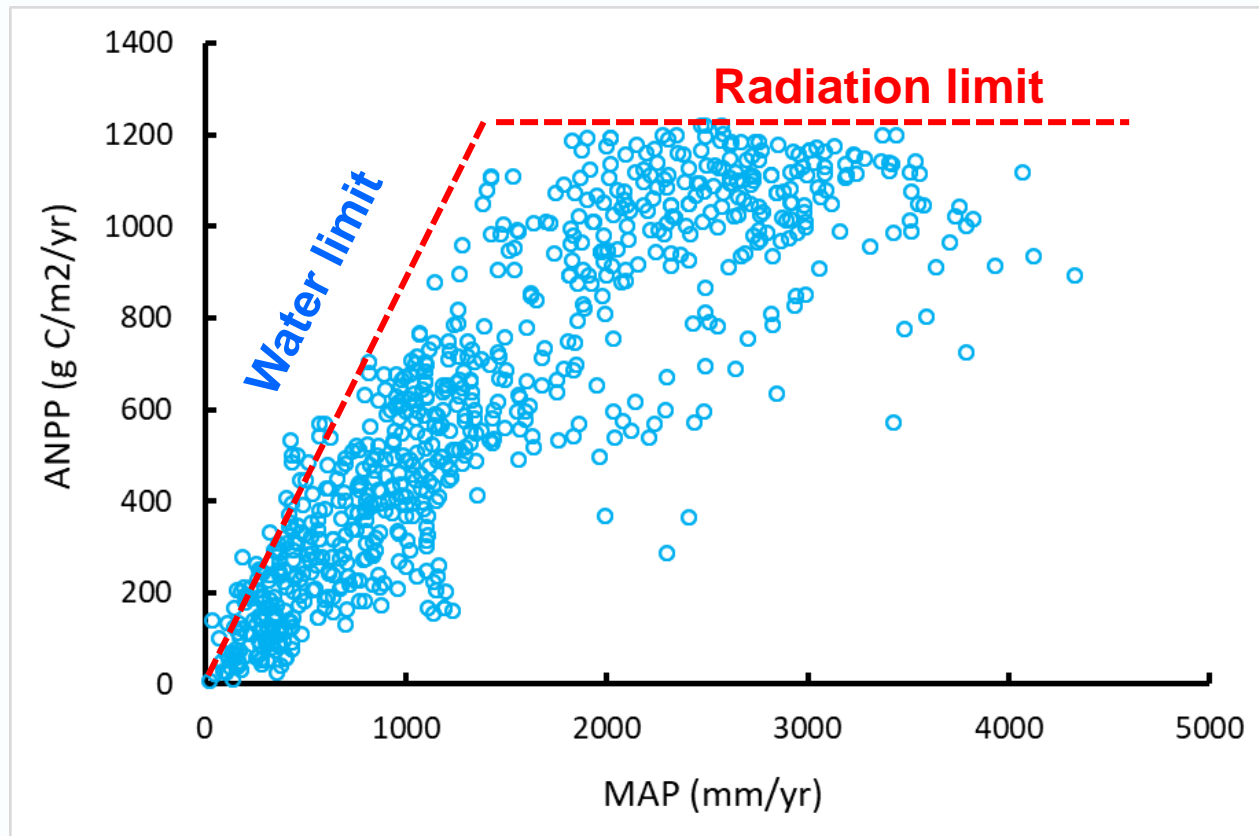
Locations of the 688 field-observed ANPP

Key controls on ANPP



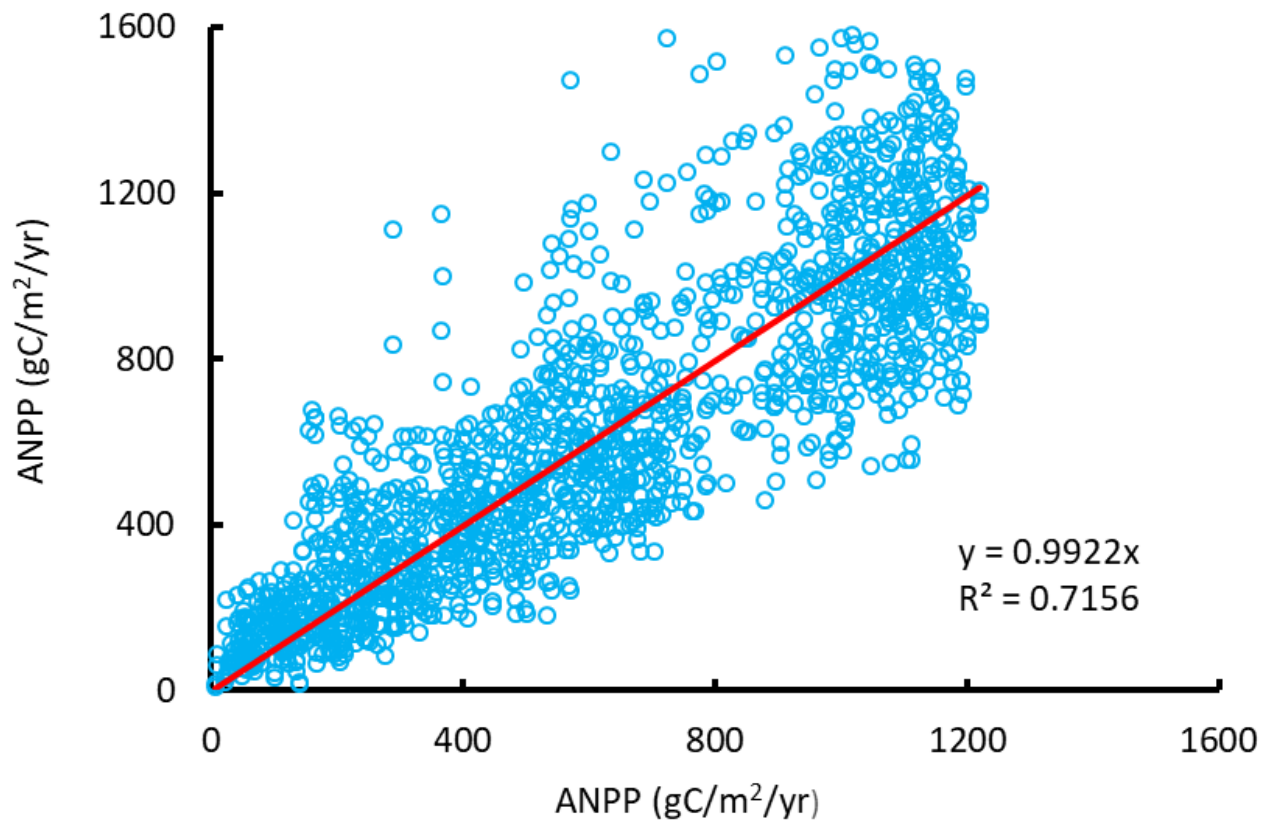
Estimation of mean annual ANPP using IRM

The Global Primary Production Data Initiative (*GPPDI*) data (1508 points)



- AMP and ANPP exhibits relationship similar to the Budyko curve

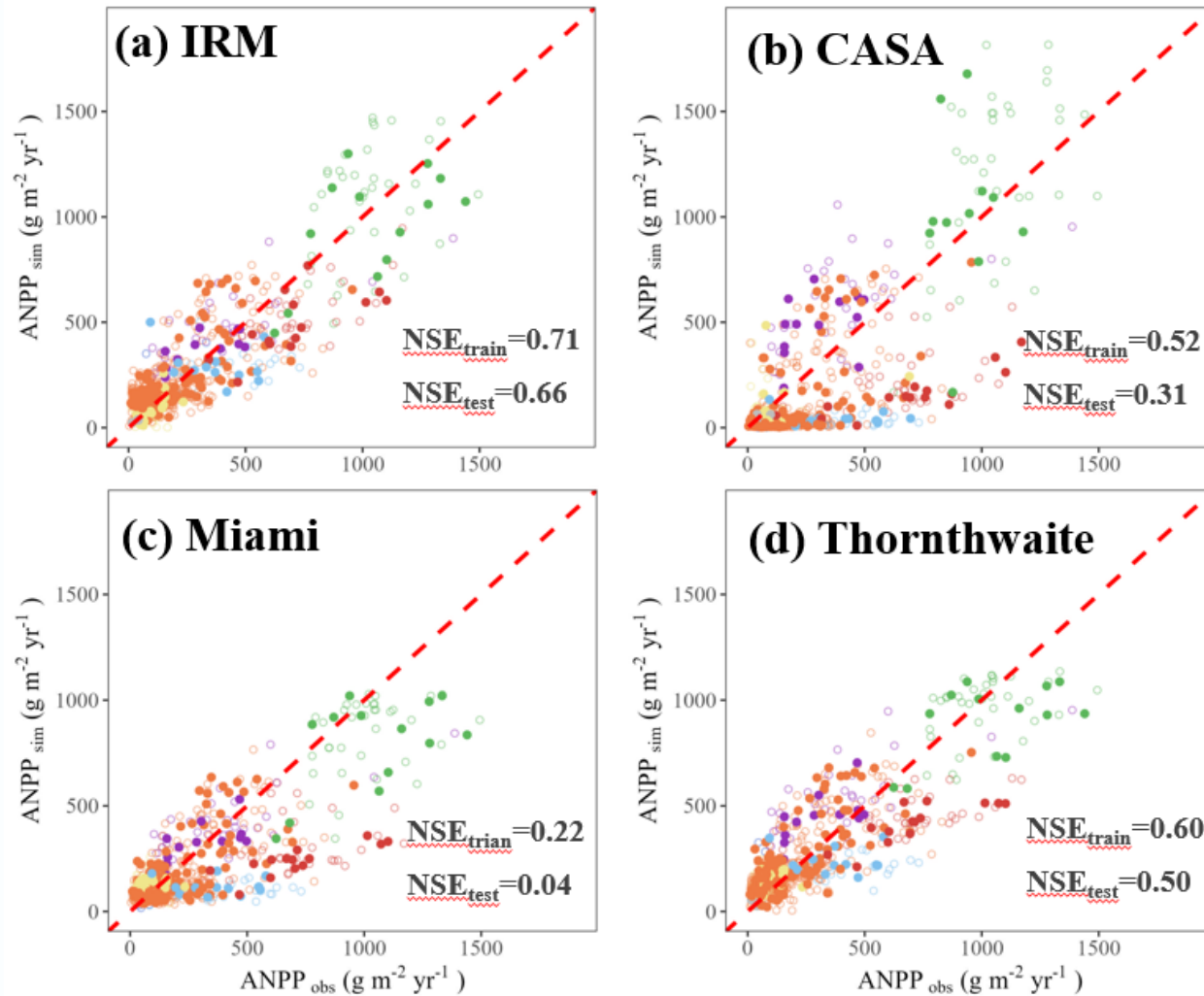
Estimation of mean annual ANPP using IRM



- IRM can provide reasonable estimates of mean annual ANPP

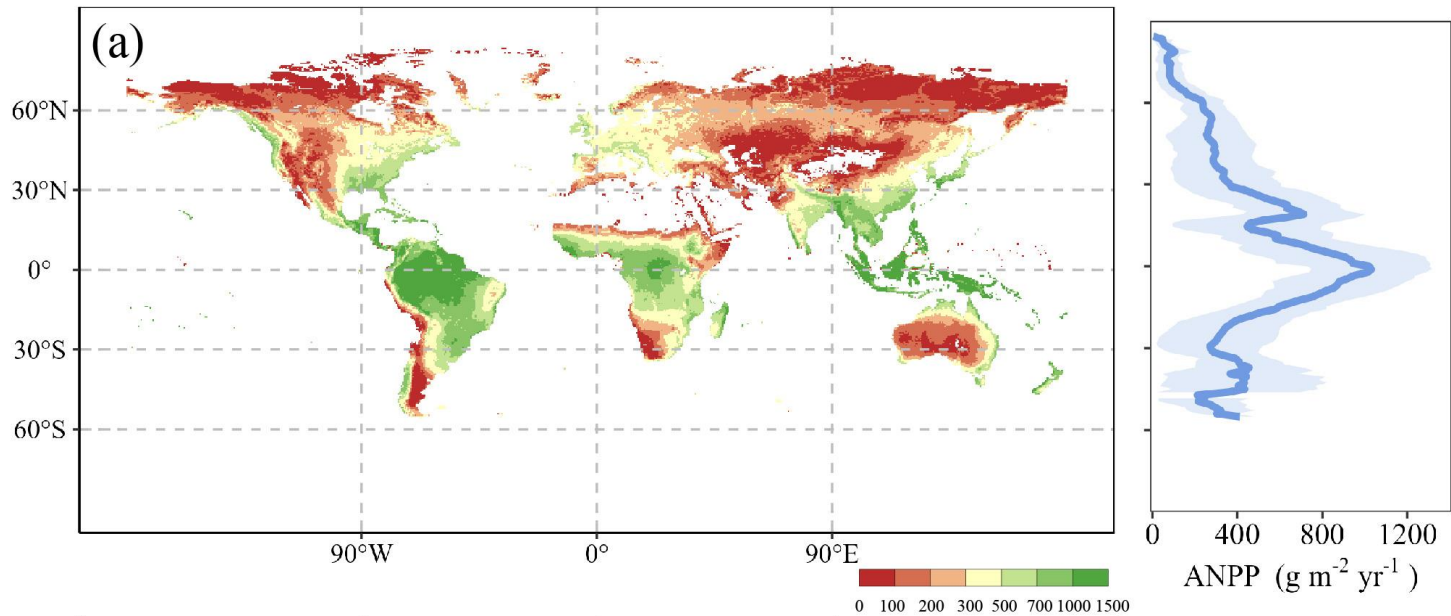
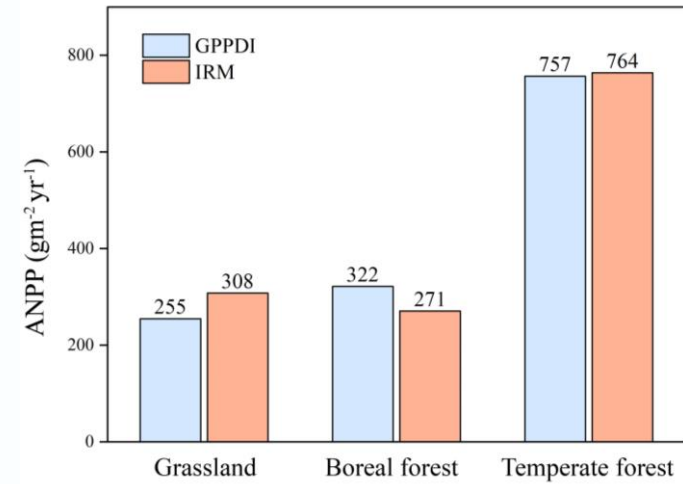


Estimation of mean annual ANPP using IRM

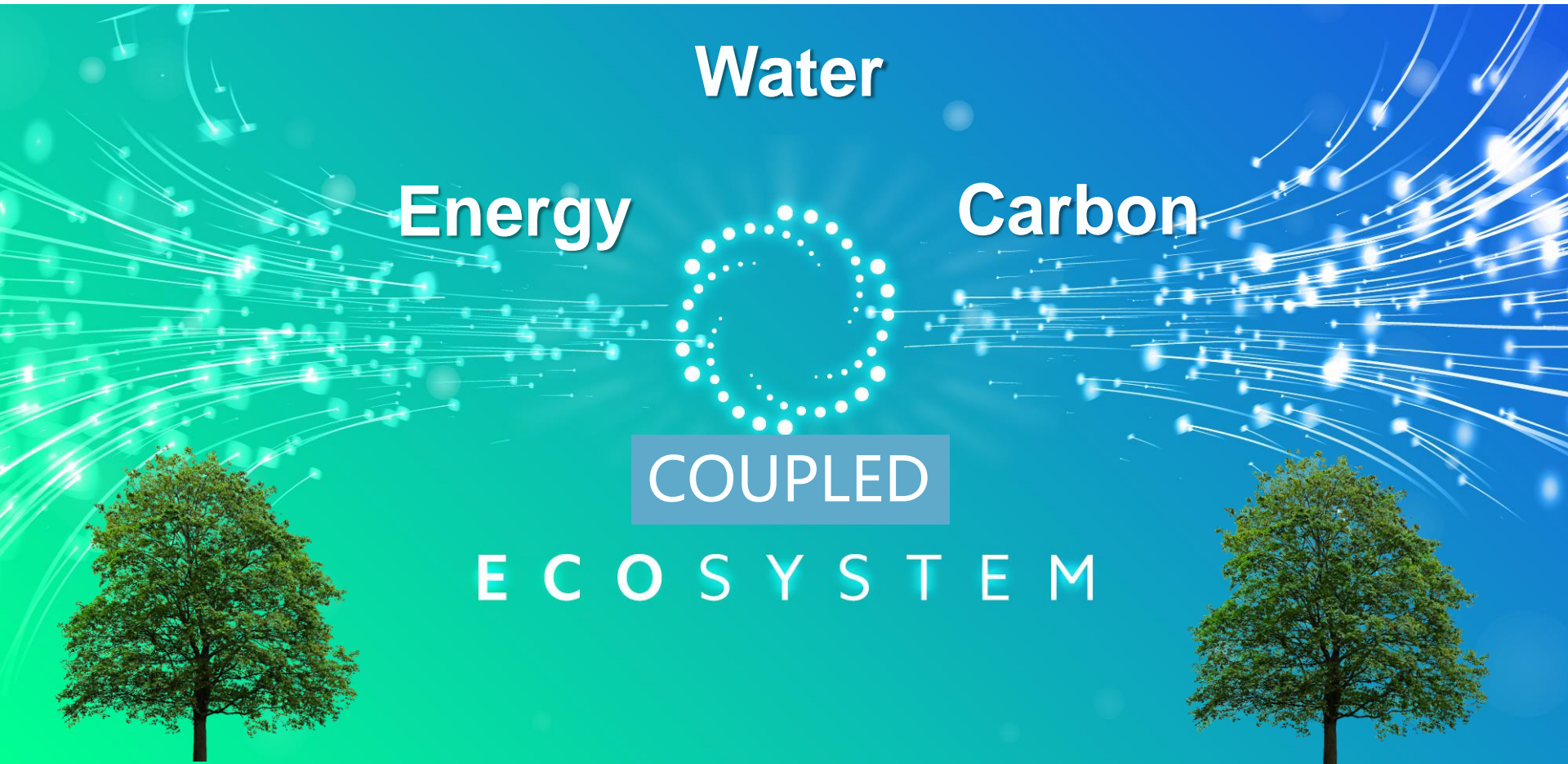


➤ IRM can provide reasonable estimates of mean annual ANPP

Estimation of global ANPP



IRM is a useful tool for modelling carbon & water fluxes



Summary

- Carbon and water coupling is important for ecohydrology and ecosystem services.
- Dominant controls of ecosystem water and carbon balance include precipitation and radiation.
- The integrated rate methodology (IRM) can be used to estimate mean annual ET and ANPP.
- Uncertainties in the ET and ANPP estimates include temperature and age effects.
- Further studies will be conducted to examine sensitivity parameters in the IRM and their effects on water and carbon modelling.

Thank you !

